The influence of the Atlantic Multidecadal Oscillation on the eastern Andes low-level jet and precipitation in South America

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## **Precipitation Changes in South America**

#### Liebmann et al. 2004:

- Significant positive trend in Jan-Mar precipitation
- (rain gauge data)





FIG. 3. (a) Climatology of monthly rainfall in southern Brazil (mm month<sup>-1</sup>). (b) JFM rainfall (in mm season<sup>-1</sup>) in southern Brazil vs year. Dashed curve represents linear, least squares fit.

#### Seager et al. 2010:

- Positive trend in annual precipitation
- Warm SST in subtropical Atlantic
- Cool tropical Atlantic (AMO<sup>-</sup>)



#### South America Monsoon System (SAMS)



Jones, C., and L. M. V. Carvalho, 2013: Climate change in the South American Monsoon System: present climate and CMIP5 projections. *Journal of Climate* 



- **1. Easterly winds**
- 2. South America low-level jet (SALLJ)
- **3. SACZ**
- 4. South Atlantic subtropical highpressure
- **5. Upper-level Bolivian High**

### South America Low-level Jet (SALLJ)

- Low-level jet: maximum winds 500-1600 m a. g.
- Max wind speeds near Bolivia ("Andes elbow")
- **Nocturnal: maximum winds 8 PM- 2AM Bolivia local time**
- Transports large amounts of moisture
- **Form mesoscale convective complexes**
- Present year round (unlike Great Plains LLJ)



## The Atlantic Multidecadal Oscillation (AMO)

<u>AMO</u>

- Positive/cold phase ⇒ warm/cold North Atlantic
- Instrumental record: 60-80 year period
- Positive: ~1924-1965
- Negative: ~1966-1996
- Precipitation changes
   Sahel, southeast US, northeast Brazil
- Natural mode of Atlantic Meridional Circulation (AMOC)
- Paleoclimate evidence
- In a few coupled models [e.g., HadCM3, Knight et al, 2006]



Several studies show significant influence of the AMO on precipitation variability in South America:

- Northeast Brazil [e.g., Kayano et al. 2014, 2016, 2018]
- Positive trend in La Plata River Basin [e.g., Seager 2010]
- Speleothem records in Botuvera, southern Brazil show 60-80 yr variations [~8000 years record; Bernal et al. 2016]
- Discharge sediments in La Plata River drainage show 60-80 yr variations [~4500 years record; Chiessi et al. 2009]



Jones, C. and L. M. V. Carvalho, 2018: The influence of the Atlantic Multidecadal Oscillation on the eastern Andes low-level jet and precipitation in South America. *Nature Clim. Atmos. Sci.* (In Press)

## Data

 European Centre for Medium-Range Weather Forecast Reanalysis of the 20<sup>th</sup> Century (ERA20C)

- 3-hourly, ~125 km
- January 1900- December 2010
- U, V at 850-hPa, 700-hPa and precipitation

 ERA20C assimilates only surface and mean sea level pressure, and surface marine wind observations

# How reliable is the ERA20C over South America?



Vertical profiles of mean wind speed (m s<sup>-1</sup>) at 06 UTC near Santa Cruz (*top*) and Mariscal (*bottom*). Profiles are averaged over all SALLJ days identified during the SALLEX field campaign for which observations are available. Radiosonde profiles are in black and reanalysis profiles are colored according to the legend.

Climatology, variability and recent changes in the SALLJ Montini et al. (2018), JGR (submitted)

#### Mean vertical wind speed profiles over Santa Cruz de la Sierra (Left) and Mariscal Estigarribia (right): January-February 2003



#### Correlation between ERA20C and GPCP precipitation during January 1979-December 2010





**Correlation with rain gauge data (1900-2010):** 

- Fortaleza 0.77
- Quixeramobim 0.75

## Methods

 European Centre for Medium-Range Weather Forecast Reanalysis of the 20<sup>th</sup> Century (ERA20C)

- 3-hourly, ~125 km
- January 1900- December 2010
- U, V at 850-hPa, 700-hPa and precipitation
- Detrend 3-hr time series
- Identify SALLJ occurrences with modified Bonner criteria:
  - Northerly winds at 850-hPa ≥ P<sup>th</sup> ms<sup>-1</sup>
  - Vertical wind shear [850-700] hPa ≥ P<sup>th</sup> ms<sup>-1</sup>
  - v-wind < 0</p>
  - v-wind > u-wind
  - Where P<sup>th</sup> are 75<sup>th</sup> seasonal percentiles
    - 6-hourly reanalysis
    - Determined for each location (Santa Cruz, Mariscal) separately
    - Determined for each season separately

## Methods

Daily SALLJ occurrences in Santa Cruz and Mariscal
 (1 Jan 1900 – 31 December 2010)

Normalized monthly SALLJ index:

$$SALLJ_{index}(t) = \left[\frac{N(t)}{CN}\right] x \left[\frac{VS(t)}{CV}\right]$$

N (t) = number of SALLJ days in the given month  $CN \equiv climatological number of SALLJ days in the month$   $VS \equiv mean wind speed (850-hPa) of the SALLJ days in the given month$  $CV \equiv climatological wind speed (850-hPa) of SALLJ days in the month$ 

Apply low-pass filter retaining variations > 8 years

AMO index from Trenberth & Shea: detrended 10-year low-pass filtered (SST) anomalies (HadISST data) over the North Atlantic basin

#### The AMO and SALLJ



## **Correlation SALLJ index and SST**



- SALLJ index: no trend, filtered > 8 years
- SST anomalies: no trend, filtered > 8 years











Consistent with model experiments during AMO<sup>-</sup> [Seager et al. 2010]

#### **AMO, SALLJ and Precipitation**



## Conclusions

Remarkable influence of the AMO on decadal-tomultidecadal variability of the SALLJ:

- Negative AMO phases 
   → negative precipitation
   anomalies over northern Amazon and Atlantic ITCZ
- Increased cross-equatorial flow & outflow over the negative precipitation anomalies => enhanced SALLJ
- Enhanced SALLJ ⇒ cyclonic circulation & enhanced precipitation over southern Brazil, Paraguay, northern Argentina
- Dynamical mechanism to explain decadal-tomultidecadal variations in precipitation (paleo records) in southeast South America

## **Tropical heating at the equator**

Some simple solutions for heat-induced tropical circulation



10N

10N



## **Mechanisms of the SALLJ**

- Deflection of trade winds crossing Amazon basin
- Purely local, topographically generated feature, driven by dry dynamics, possibly modified by moist convection on the Andean slopes
- externally forced feature, produced by variations in the pressure field in northern Argentina: cross Andes flow from Pacific
- Propagation of low-level wind bursts from the North Atlantic toward the La Plata basin through the Amazon basin

Vera et al. (2006)

#### **Theories for the Great Plains LLJ (nocturnal)**

Heating/cooling over sloping terrain (Holton 1967): Diurnal heating cycle on the sloping Great Plains generates vorticity baroclinically that leads to diurnal oscillation in BL winds

> Horizontal temperature gradients



**FIGURE 3.8** Relationship between vertical shear of the geostrophic wind and horizontal thickness gradients. (Note that  $\delta p < 0$ .)

Vertical shear in geostrophic winds

$$\begin{bmatrix} \frac{\partial u_g}{\partial z} = -\frac{g}{fT} \frac{\partial T}{\partial y} \\ \frac{\partial v_g}{\partial z} = \frac{g}{fT} \frac{\partial T}{\partial x} \end{bmatrix}$$

Shapiro, A., et al., 2016: A Unified Theory for the Great Plains Nocturnal Low-Level Jet. J. Atmos. Sciences, 73, 3037-3057.

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## **Theories for the Great Plains LLJ (nocturnal)**

 Mountain-blocking (Wexler 1961): Blocking of easterly trade winds by the Rocky Mountains yields strong southerly jet
 Inertial Oscillation (Blackadar 1957): Sudden decrease of frictional stress in the boundary layer (BL) around sunset destroys Ekman balance and accelerates air parcels.



Shapiro, A., et al., 2016: A Unified Theory for the Great Plains Nocturnal Low-Level Jet. *J. Atmos. Sciences*, 73, 3037-3057.

## **Climate Variability and Change in South America**

How will climate change in South America in the coming decades?



How does natural climate variability on decadal-tomultidecadal scales work in South America?

How strong is the anthropogenic climate forcing relative to natural variability?









# South America Low-level Jet (SALLJ)







Marengo, J. A., et al., 2004: Climatology of the low-level jet east of the Andes as derived from the NCEP-NCAR reanalyses: Characteristics and temporal variability. J. Climate, 17, 2261-2280.